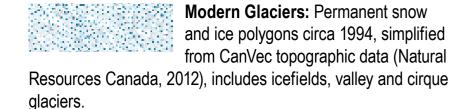
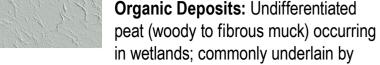


#### Holocene





# Alberta Geological Survey Map 601 Legend Surficial Geology of Alberta



fine-grained, poorly drained glaciolacustrine or lacustrine deposits; in some places underlain by till; includes marshes, swamps, bogs, and fens.



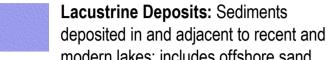
Colluvial Deposits: Sediments that have reached their present position as a result of direct, gravity-induced

movement; commonly occurs as slope and slump deposits confined to valley sides and floors; includes bedrock and surficial materials; in places, includes a significant component of fluvial deposits, as these two units are inseparable at this map scale.



by streams and rivers; synonymous with alluvium; includes poorly to well-

sorted, stratified to massive sand, gravel, silt, clay, and organic sediments occurring in channel and overbank deposits; in places, includes a significant component of colluvial deposits as these two units are inseparable at this map scale.



modern lakes; includes offshore sand, silt, and clay, minor organic deposits; may also include minor littoral (nearshore) beaches and bars composed of sand, silt, and minor gravel.

Eolian Deposits: Wind-deposited

sediments comprising well-sorted,



# Introduction

Map 601 portrays a generalized compilation of the surficial geology of Alberta at 1:1 000 000 scale using previously published maps from the Alberta Geological Survey, the Geological Survey of Canada, and Environment Canada, as well as university theses and new data. If detailed information is required on the surficial geology of specific areas, the original source maps (Figure 1) should be consulted.

#### Data Sources

The Surficial Geology Map of Alberta is based on an amalgamation of 65 previously published surficial geology maps that cover about 90% of the province (Figure 1). These maps date back to the 1960s, some of which used unstandardized legends and were published at scales ranging from 1:50 000 to 1:500 000. These maps were digitized using ArcGIS and unit polygons were reclassified according to the Map 601 legend to produce a seamless mosaic of the surficial geology of Alberta. The remaining unmapped areas of Alberta were completed by field-tested Light Detection and Ranging (LiDAR) and airphoto interpretations and predictive mapping at 1:500 000 to 1:1 000 000 scale as part of this compilation (Figure 2) and were incorporated into the seamless provincial mosaic.

Boundary discrepancies between units in adjacent map areas were largely resolved during the reclassification process. However, in areas with significant discrepancies, unit-specific reclassification was undertaken along polygon boundaries. Furthermore, if original maps included symbology rather than unit polygons, then these symbols were used to reclassify units according to the standard legend and incorporated into Map 601. Further unit reclassification included data from other published sources, such as in areas where fluting was not recognized on an original map but was subsequently detected on digital elevation models (e.g., Ó Cofaigh et al., 2010).

The corrected seamless provincial mosaic was then generalized using GeoScaler software published by the Geological Survey of Canada (Huot-Vézina et al., 2009). The resulting simplified surficial geology layer is shown overlaid on a hill-shaded Shuttle Radar Topography Mission digital elevation model (U.S. Geological Survey, 2000) to convey the effects of topography.

Due to the compilation of surficial geology maps spanning five decades and published at a range of scales, base features shown on this map do not always conform to the generalized geological unit boundaries, since these features were derived from the most up-to-date 1:1 000 000 scale digital data (Natural Resources Canada, 2009). Consequently, there may be areas on the map that appear inconsistent; for example, where modern rivers flow outside of fluvial sediment polygons. Nonetheless, at this map scale, the base features and the geological units have a positional accuracy of ± 1.5 km.

## Surficial Geology Map of Alberta

Map 601 illustrates the geology at or near the modern land surface (to a depth of approximately 2 m) The relative abundance of the map units is undivided moraine (24%), fluted moraine (7%), stagnant ice moraine (19%), and ice-thrust moraine (2%), as well as glaciolacustrine (18%), organic (10%), colluvial (5%), fluvial (4%), glaciofluvial (4%), eolian (3%), lacustrine (1%), and preglacial fluvial (0.1%) deposits. Outcrops of bedrock comprise approximately 3% of the surficial geology of Alberta and primarily occur within the Rocky Mountains and in the Canadian Shield adjacent to Lake Athabasca (Pană et al., 2013; Prior et al., 2013).

glaciolacustrine sediments. These were deposited in glacial lakes that ponded against the retreating ice margins, many of which extended distally along the axes of a number of major river valleys in Alberta. The most regionally extensive distribution of glaciolacustrine deposits occurs across the Northern Alberta Lowlands (Figure 5), extending from Lesser Slave Lake, northwards along the flanks of the Peace and Hay rivers towards the provincial boundary to the north and west, and Lake Athabasca to the east.

Glaciofluvial deposits (4%) represent a minor surficial geological unit and are associated with the deglaciation of Alberta during the late Pleistocene. These deposits generally form linear or irregularly shaped bodies and are found primarily within two major terrain types. The first is lowlands incised by major glacial meltwater channels, which include the Hay and Athabasca rivers in northwestern and northeastern Alberta, respectively. In these settings, glaciofluvial deposits occur along the margins of these channels. The second is along the flanks of stagnant ice moraine-dominated uplands, including the Buffalo Head Hills and the Caribou, Birch, and Stony mountains (Figure 6). However, the most extensive area of glaciofluvial deposits occurs between Lake Athabasca and Fort McMurray, near the sandstone-rich Athabasca Group (Slimmon and Pană, 2010).

#### Pleistocene to Holocene

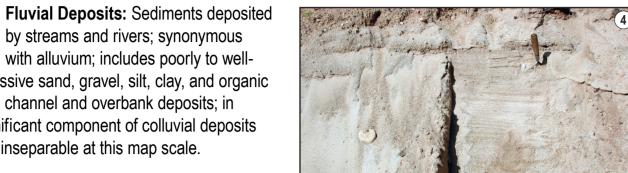
Eolian deposits (3%) range in age from late Pleistocene to Holocene (Wolfe et al., 2004). Many were deposited immediately following deglaciation when newly exposed sediments were not yet stabilized by vegetation. Consequently, eolian deposits are commonly located adjacent to glaciofluvial units, as well as modern fluvial terraces and floodplains due to their derivation from redeposited fine-grained sediments. These deposits are particularly prevalent along the axes of the Athabasca and Peace rivers, where they occur within a succession of locally extensive dune fields. Eolian deposits also occur across areas of sandy moraine in northeastern Alberta.

#### Holocene

The most widespread, postglacial, surficial geological unit comprises organic deposits (10%), which occur throughout northern Alberta, and form a significant component of the boreal forest. Organic deposits are common across the lowlands between the Buffalo Head Hills, Birch Mountains, and Stony Mountain (Figure 6), where they typically overlie and are interspersed with glaciolacustrine deposits. Organic deposits are also associated with areas of poorly drained moraine that drape parts of the Caribou and Birch mountains and the Cameron Hills. At these elevations, many organic deposits contain discontinuous areas of permafrost.

Minor postglacial, surficial geological units across the Alberta landscape include colluvial (5%), fluvial (4%), and lacustrine (1%) deposits. Colluvium forms a locally significant unit fringing the slopes of the Rocky Mountains and Foothills, as well as on the flanks of the Swan Hills, and the Birch and Caribou mountains. Colluvial deposits are also prominent along the margins of major valley systems in Alberta. Due to its association with river valleys, colluvium and fluvial deposits, which are deposited adjacent to modern rivers, are often inseparable at the scale of this map. The ability to differentiate these two types of deposits within a particular valley segment depends on the relative proportion of the valley slopes and the floodplain.





#### **Pleistocene to Holocene**



medium- to fine-grained sand and minor silt; generally massive to locally cross-bedded or ripple laminated; includes both active and vegetated dunes and sand sheets

#### Pleistocene



Glaciolacustrine Deposits: Sediments deposited in or along the margins of glacial lakes; includes a) offshore

sediment; rhythmically laminated to massive fine sand, silt, and clay, locally containing debris released by the melting of floating ice; and b) littoral (nearshore) sediments; massive to stratified, well-sorted silty sand, pebbly sand, and minor gravel; occurs in beaches, bars, and deltas.



**Glaciofluvial Deposits:** Sediments deposited by glacial meltwater in subaerial, subaqueous, and subglacial

environments; sediment ranges from massive to stratified, poor to well sorted, coarse to fine grained; in places, this unit includes till; may show evidence of ice melting (slumped structures).



Moraine: Diamicton (till) deposited directly by glacial ice with a mixture of clay, silt, and sand, as well as minor

pebbles, cobbles, and boulders; characterized by a lack of distinctive topography. Locally, this unit may contain blocks of bedrock, stratified sediment, or lenses of glaciolacustrine and/or glaciofluvial sediment.



Fluted Moraine: Glacially streamlined sediments, mainly till; terrain varies from alternating furrows and ridges to

elongated smoothed hills which parallel the inferred local ice-flow direction; includes flutes, drumlins, and drumlinoids



The age, origin, thickness, and distribution of the sedimentary succession overlying the bedrock surface in Alberta varies greatly, with thickness ranging from less than 1 m to as much as 445 m (Figures 3 and 4). These sediments document geological processes that were active across the Alberta landscape during the Cenozoic, the most recent geological era, and include Paleogene to Neogene fluvial deposits and glaciogenic materials deposited during Quaternary glaciation, as well as postglacial deposits. The general relationship between the bedrock topography and the origin and distribution of surficial deposits shows that large parts of the modern Alberta landscape relate to the topography of the underlying bedrock.

## Sediment Distribution

#### Paleogene to Neogene

Preglacial fluvial deposits occupy isolated, gravel-capped uplands within widely dispersed settings across the Interior Plains of Alberta, including Cypress Hills, Hand Hills, Swan Hills, and Marten Mountain (the western portion of the Pelican Mountains; Figures 5 and 6). The preglacial fluvial deposits occur within peneplains that represent the erosional remnants of eastward-flowing fluvial systems formed during the Paleogene to Neogene following regional uplift associated with the Laramide Orogeny (Leckie, 1989, 2006; Leckie and Cheel, 1989; Leckie et al., 2004). During ongoing uplift, eastwardflowing rivers continued to erode these peneplains, redepositing gravel at successively lower elevations across the landscape.

#### Pleistocene

Map 601 reveals that the most abundant geological units that cover the Alberta landscape originated during the Pleistocene, as a result of the advance and retreat of Quaternary ice sheets that deposited moraine (52%) and glaciolacustrine deposits (18%).

Ice advance was responsible for the widespread deposition of moraine extending from plains in the southern, central, and northeastern parts of Alberta, to uplands in the west and north (Figure 5). During the last glaciation (culminating approximately 22 000 years ago; Mix et al., 2001; Dyke et al., 2002), the Laurentide Ice Sheet advanced south to southwestwards across Alberta, against the regional topographic gradient, and converged with the Cordilleran Ice Sheet, which was advancing eastwards onto the Southern Alberta Plains from the Rocky Mountains (Figure 4).

Moraine is classified into four units, in descending order of its spatial distribution across the province: undivided, stagnant ice, fluted, and ice thrust.

Undivided moraine is widespread across uplands in northern Alberta, including the Cameron Hills, Caribou Mountains, the northern flank of the Clear Hills, and the eastern flank of the Birch Mountains (Figure 6). Undivided moraine also extends across parts of the Northern Alberta Lowlands, adjacent to the flanks of the Caribou and Birch mountains. It is also common across the Saskatchewan Plains, adjacent to the confluence of the Athabasca and Clearwater rivers, as well as across uplands and benchlands that parallel the Rocky Mountain Foothills in western and southern Alberta.

Stagnant ice moraine is a widespread unit in north-central Alberta, particularly across the Buffalo Head Hills and the southern flanks of the Birch Mountains, Clear Hills, and Swan Hills (Figure 6). It is also abundant across uplands within the western and eastern Alberta plains, particularly along the flanks of three prominent, low-relief corridors of fluted moraine extending south between Edmonton and Calgary, to the east and west of Red Deer, and southeast from Lac La Biche. The central and western corridors of fluted moraine terminate at a broad arc, originally termed the Lethbridge Moraine (Stalker, 1977), which extends from Strathmore to Medicine Hat and contains a suite of glacial landforms, including ice-thrust moraine. Ice-thrust moraine occurs in other places in Alberta, most commonly southeast of Lac La Biche, on the northern slopes of the Pelican Mountains and the Swan Hills, to the southwest of the Birch Mountains, and adjacent to the southern margins of many lakes in central Alberta.

During deglaciation of Alberta (initiating approximately 14 500 years ago; Dyke, 2004), the Laurentide Ice Sheet retreated northwards and eastwards into the Northwest Territories and Saskatchewan, and the Cordilleran Ice Sheet retreated westwards into the Rocky Mountains. The most widespread surficial geological unit associated with the retreat of the Laurentide Ice Sheet comprises late Pleistocene

74M

Recently deposited lacustrine deposits are rare in Alberta (1%), the most extensive areas occurring around the margins of Zama and Lesser Slave lakes.

#### Significance to the History and Development of Alberta

The significance of fluvial, glacial, and postglacial processes and their associated deposits is evident across the modern Alberta landscape. The cumulative effect of these processes during the Cenozoic history of Alberta has strongly influenced the topography, hydrology, hydrogeology, pedology, and ecology of the province and has an important role in modern land-use patterns and development.

The location of many of Alberta's major rivers, including the North and South Saskatchewan, Athabasca, Peace, and Hay rivers, in part relate to the courses of ancestral Paleogene to Neogene river systems. Some paleovalley systems associated with these ancestral rivers are filled by a substantial thickness of Cenozoic sediment. In places, this sediment forms regionally significant near-surface aquifers (Figure 4; Barker et al., 2011). In some areas of Alberta, the continued evolution of these rivers during the Quaternary was influenced by the location of glacial meltwater channels that either drained glacial lakes or formed along successive ice-marginal positions as ice sheets retreated across the landscape. For example, in southern and eastern Alberta, many rivers follow courses oblique to the regional topographic gradient due to their deflection around broad lobes of retreating Laurentide ice that extended southwards into the drainage basins of these rivers during regional deglaciation (Klassen, 1989).

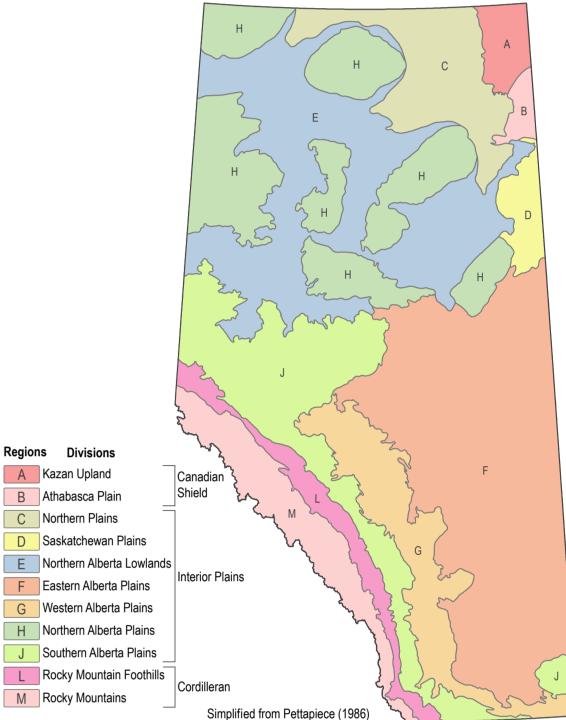
In addition to insights into the geological history of Alberta that these river courses provide, they also played a significant role in the historical development of the province. During the late 18th and early 19th centuries, Alberta's rivers were used as 'fluvial highways' that enabled access to much of the Interior Plains and Cordillera (Scheelar and Macyk, 1972). Because of the historical settlement of these river valleys, many of Alberta's cities now occupy the basins of former glacial lakes that were impounded by the Laurentide Ice Sheet as it retreated from the drainage basins of these rivers, including Medicine Hat, Lethbridge, Calgary, Red Deer, Edmonton, Peace River, Grande Prairie, and Fort McMurray.

Alberta's major soil zones and natural regions (Downing and Pettapiece, 2006) are significantly influenced by the underlying geology. For example, wetlands are preferentially developed in areas underlain by low permeability sediments, primarily glaciolacustrine deposits. Consequently, large areas of the boreal forest in the Northern Alberta Lowlands and Northern Plains (Figure 5) contain organic deposits that accumulated within the lowland basins of former glacial lakes. Glaciolacustrine deposits are typically associated with low-relief terrain, comprising stone-free, fine-grained sediment. Soils derived from this material are consequently more suitable to agriculture, as can be seen by the widespread distribution of arable land on the lowlands adjacent to the Peace River in northwest Alberta and along the North Saskatchewan River in the Edmonton region of central Alberta.

In contrast, areas of moraine, which comprise a heterogeneous mixture of rock fragments within a silty to sandy matrix, were historically more challenging for agriculture because farmers had to clear the cobbles and boulders from the surface to reduce damage to farm equipment during tillage. Furthermore, because some types of moraine are typically associated with rolling to higher relief topography, those landscape elements were generally less conducive to efficient agriculture prior to the development of modern farming technology. In addition to agriculture, areas of moraine are also associated with a range of other land uses, including grazing and forestry. Surficial deposits dominated by coarse-grained sediments form a minor component of the Alberta landscape. However, in northeast Alberta, extensive eolian and glaciofluvial deposits across the Athabasca Plain have a strong ecological influence on the landscape and contribute to the development of unique vegetation communities and Jack pinedominated uplands (Downing and Pettapiece, 2006).

The underlying geology also influences postglacial processes, and in some areas of Alberta, these processes may have a significant effect on modern infrastructure development. For example, the owlands flanking the Peace River are one of the most historically active mass-movement areas in Western Canada. This mass-movement activity is influenced by the distribution of glaciolacustrine deposits in the area, particularly where thick deposits of glaciolacustrine sediment infill preglacial paleovalleys, which are prone to slope instabilities, particularly when undercut by present-day rivers. Infrastructure associated with urban development in these areas, such as the town of Peace River, has been, and continues to be, affected by these instabilities (Morgan et al., 2012).

#### Figure 6. Surface Topography and Photo Locations





Regions

resulting from the collapse and slumping of englacial and supraglacial debris due to the melting of buried stagnant ice at the glacier margin; sediment is mainly till but locally includes stratified glaciolacustrine or glaciofluvial sediments; characterized by low- to high-relief hummocky topography.

Stagnant Ice Moraine: Sediments









Ice-Thrust Moraine: Material formed from the glaciotectonic displacement of blocks or rafts in a more or less intact

state; may include syngenetic till and masses of pre-existing sediments and/or bedrock; characterized by high- to moderaterelief glaciotectonic moraines; includes hill-hole pairs, rubble moraine, and thrust-block moraine.



# Paleogene to Neogene

Preglacial Fluvial Deposits: Sediments transported and deposited by streams and rivers prior to glaciation; Cordilleran source; predominantly well-sorted quartzite and chert gravel and cobbles with minor sand; unit generally forms erosional remnants capping isolated upland and midland peneplains.

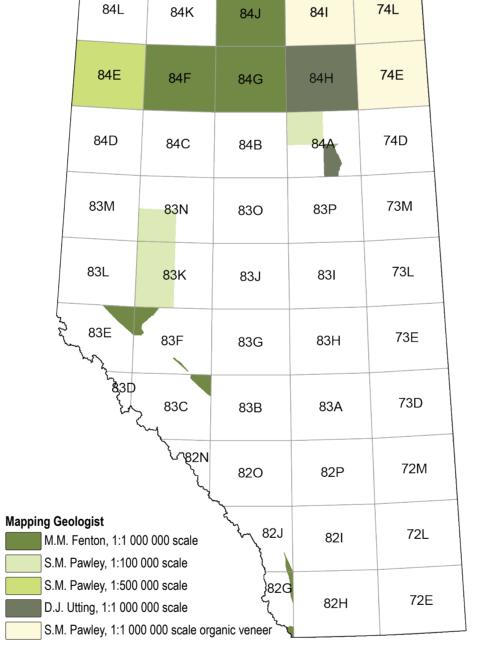


# Precambrian to Cretaceous

northeastern Alberta (igneous and minor metamorphic rock), the foothills (clastic rocks including conglomerate, sandstone, siltstone,

and shale), and in the mountains (clastic rocks, limestone, and dolostone); minor exposure on the plains (clastic rocks).

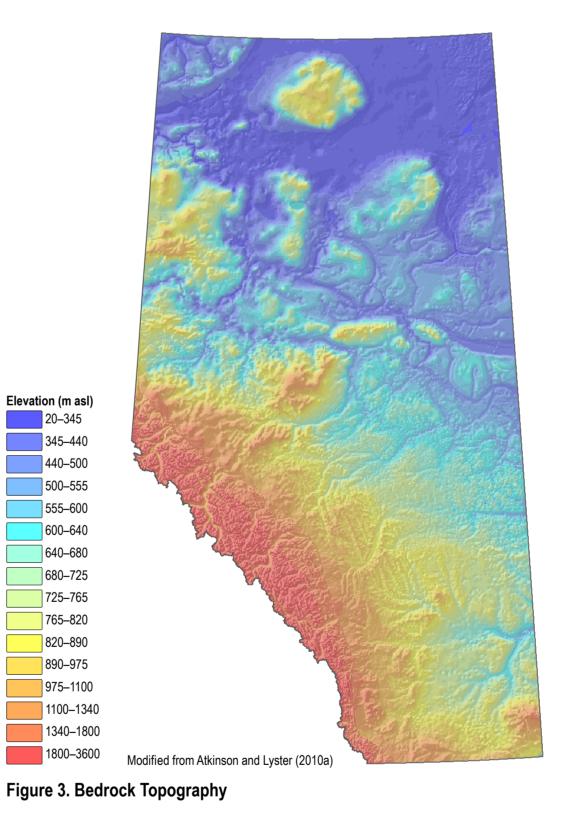


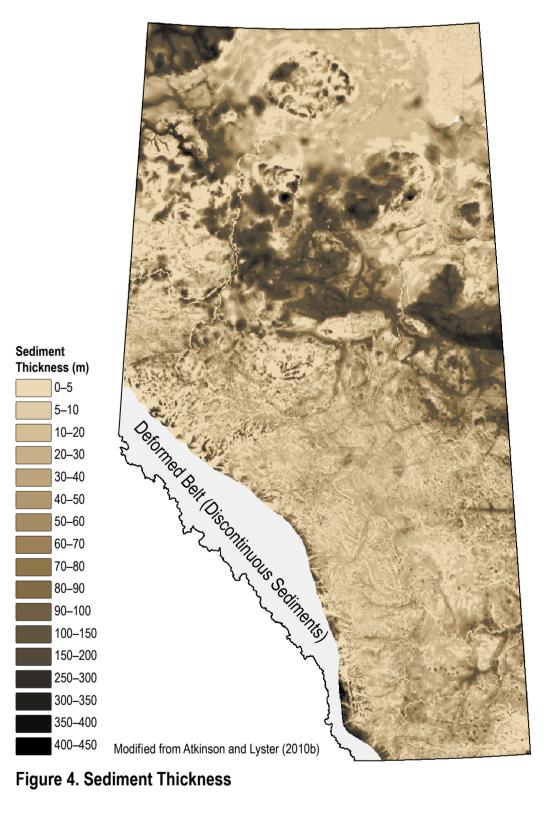


84O

84M

Figure 2. New AGS Mapping for this Compilation





**References for Compiled Data Sources (See Figure 1)** 

Bedrock: Mapped primarily in

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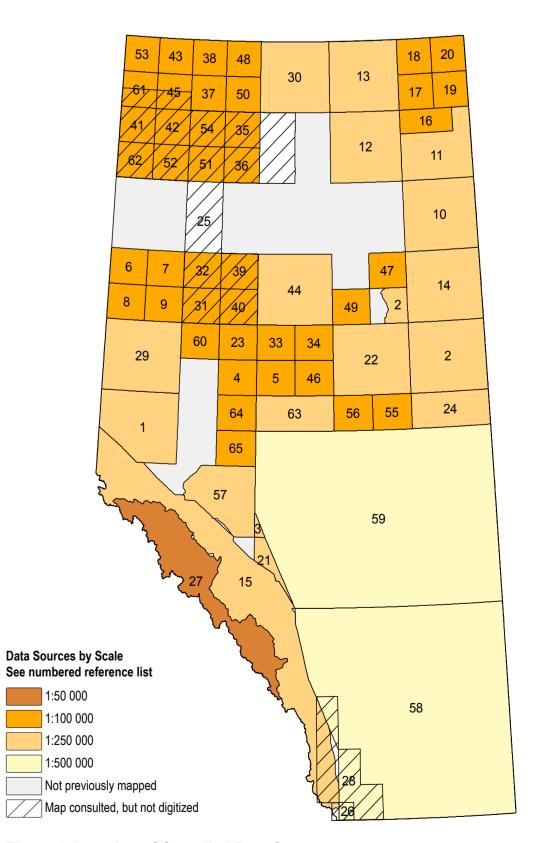
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#### **Recommended Reference Format**

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# Figure 1. Location of Compiled Data Sources

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